

# Derivation of Hydroperoxyl Radical Levels at an Urban Site via Measurement of Pernitric Acid by Iodide Chemical Ionization Mass Spectrometry

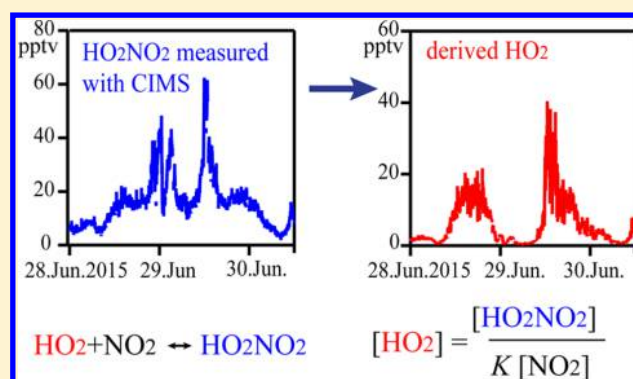
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## Supporting Information

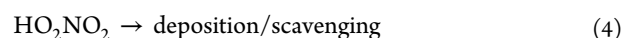
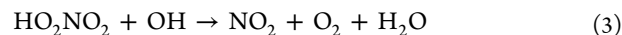
**ABSTRACT:** Hydroperoxyl radical ( $\text{HO}_2$ ) is a key species to atmospheric chemistry. At warm temperatures, the  $\text{HO}_2$  and  $\text{NO}_2$  come to a rapid steady state with pernitric acid ( $\text{HO}_2\text{NO}_2$ ). This paper presents the derivation of  $\text{HO}_2$  from observations of  $\text{HO}_2\text{NO}_2$  and  $\text{NO}_2$  in metropolitan Atlanta, US, in winter 2014 and summer 2015.  $\text{HO}_2$  was observed to have a diurnal cycle with morning concentrations suppressed by high  $\text{NO}$  from the traffic. At night, derived  $\text{HO}_2$  levels were nonzero and exhibited correlations with  $\text{O}_3$  and  $\text{NO}_3$ , consistent with previous studies that ozonolysis and oxidation by  $\text{NO}_3$  are sources of nighttime  $\text{HO}_2$ . Measured and model calculated  $\text{HO}_2$  were in reasonable agreement: Without the constraint of measured  $\text{HO}_2\text{NO}_2$ , the model reproduced  $\text{HO}_2$  with a model-to-observed ratio (M/O) of 1.27 ( $r = 0.54$ ) for winter, 2014, and 0.70 ( $r = 0.80$ ) for summer, 2015. Adding measured  $\text{HO}_2\text{NO}_2$  as a constraint, the model predicted  $\text{HO}_2$  with M/O = 1.13 ( $r = 0.77$ ) for winter 2014 and 0.90 ( $r = 0.97$ ) for summer 2015. These results demonstrate the feasibility of deriving  $\text{HO}_2$  from  $\text{HO}_2\text{NO}_2$  measurements in warm regions where  $\text{HO}_2\text{NO}_2$  has a short lifetime.



## INTRODUCTION

Hydroperoxyl radical ( $\text{HO}_2$ ) plays a key role in the atmosphere and is the dominant form of  $\text{HO}_x$  ( $\text{OH} + \text{HO}_2$ ). Sources of  $\text{HO}_2$  include the oxidation of CO and volatile organic compounds (VOC) by hydroxyl radicals ( $\text{OH}$ ), the photolysis of carbonyl compounds, the ozonolysis of alkenes, and the oxidation of VOC by the nitrate radical ( $\text{NO}_3$ ).<sup>1–6</sup> The last two are important sources of  $\text{HO}_2$  at night.  $\text{HO}_2$  is involved in many important atmospheric processes such as the degradation of pollutants, ozone production, and secondary organic aerosol (SOA) formation.<sup>7–9</sup> Observations of  $\text{HO}_2$  are important for understanding and constraining atmospheric chemistry.

Pernitric acid ( $\text{HO}_2\text{NO}_2$ ) is closely coupled to  $\text{HO}_2$  levels in many regions of the atmosphere.  $\text{HO}_2\text{NO}_2$  has a relatively simple chemistry in the atmosphere (Reactions 1–4).<sup>10–14</sup>



The thermal decomposition rate  $k_{-1}$  is a strong function of temperature, e.g.,  $\sim 5 \times 10^{-6} \text{ s}^{-1}$  at  $-30^\circ \text{C}$  and  $\sim 1 \times 10^{-1} \text{ s}^{-1}$  at  $30^\circ \text{C}$ .<sup>15</sup> The important role of  $\text{HO}_2\text{NO}_2$  in  $\text{HO}_x$  and  $\text{NO}_x$  ( $\text{NO} + \text{NO}_2$ ) chemistry in cold regions has been widely discussed.<sup>14,16–23</sup> In warm regions, as its thermal decomposition is fast and dominates the loss,  $\text{HO}_2\text{NO}_2$  chemistry is reduced to a rapidly established steady state with  $\text{HO}_2$  and  $\text{NO}_2$  (Reaction 1). Given proper averaging time that allows the equilibrium to be approached,  $\text{HO}_2$  can be derived from  $\text{HO}_2\text{NO}_2$  and  $\text{NO}_2$  observations.<sup>24</sup>

$$[\text{HO}_2] = \frac{k_{-1}[\text{HO}_2\text{NO}_2]}{k_1[\text{NO}_2]} = \frac{[\text{HO}_2\text{NO}_2]}{K[\text{NO}_2]} \quad (5)$$

where  $K = k_1/k_{-1}$  is the equilibrium constant and  $K$  is a function of temperature alone.<sup>15</sup>

In this work, we employ chemical ionization mass spectrometry (CIMS) to measure  $\text{HO}_2\text{NO}_2$  in metropolitan Atlanta and determine  $\text{HO}_2$  from observations of  $\text{HO}_2\text{NO}_2$  and  $\text{NO}_2$ . Hydrated iodide ( $\text{I}^-(\text{H}_2\text{O})_n$ ) is used as the reagent ion(s) for CIMS detection of  $\text{HO}_2\text{NO}_2$ .<sup>25–27</sup> The  $\text{I}^-(\text{H}_2\text{O})_n$  CIMS

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Table 1. Summary of the Measurements in Two Seasons

	species	technique	LOD	integration Time	uncertainty
winter (2014)	HO <sub>2</sub> NO <sub>2</sub>	CIMS	0.2 pptv	1 min	±22%
	N <sub>2</sub> O <sub>5</sub>	CIMS <sup>33</sup>	0.2 pptv	1 min	±25%
	NO	chemiluminescence <sup>69</sup>	4.5 pptv	1 min	±10%
	NO <sub>2</sub>	CAPS <sup>47</sup>	<60 pptv	5 s	<5%
	O <sub>3</sub>	Thermo 49C	<1 ppbv	10 s	±1 ppbv
	CO	Thermo 48C	<1 ppbv	10 s	±5%
	met. parameters <sup>a</sup>	Vantage Pro2		1 min	
summer (2015)	HO <sub>2</sub> NO <sub>2</sub>	CIMS <sup>27</sup>	0.6 pptv	1 min	±16%
	NO	Teledyne 200EU	50 pptv	1 min	±5%
	NO <sub>2</sub>	CAPS <sup>33</sup>	<60 pptv	5 s	<5%
	O <sub>3</sub>	Teledyne T400	400 pptv	10 s	±5%
	CO	IR adsorption cell	<1 ppbv	10 s	±5%
		met. parameters <sup>a</sup>	Vantage Pro2		1 min

<sup>a</sup>Meteorological parameters include *T*, *p*, humidity, wind speed, wind gust, wind direction, total solar radiation, solar UV radiation, cloud base, and precipitation rate.

technique is versatile and suitable for detections of a host of species, such as Cl<sub>2</sub>, Br<sub>2</sub>, ClO, BrO, HOCl, HOBr, N<sub>2</sub>O<sub>5</sub>, HNO<sub>3</sub>, HONO, ClNO<sub>2</sub>, HCOOH, HNCO, and CH<sub>3</sub>C(O)-OOH.<sup>28–39</sup> Consequently, HO<sub>2</sub>NO<sub>2</sub> observations and determination of HO<sub>2</sub> can be simultaneously performed with measurements of other species using the same instrument. The goal of this work is to evaluate the feasibility of deriving HO<sub>2</sub> from HO<sub>2</sub>NO<sub>2</sub> observations in a warm polluted environment. This work on deriving HO<sub>2</sub> from HO<sub>2</sub>NO<sub>2</sub> observations utilizing I<sup>−</sup>(H<sub>2</sub>O)<sub>*n*</sub> as a reagent ion was also motivated as we have found that we cannot directly measure HO<sub>2</sub> with I<sup>−</sup>(H<sub>2</sub>O)<sub>*n*</sub> without significant interferences as detailed by Sanchez et al.<sup>40</sup>

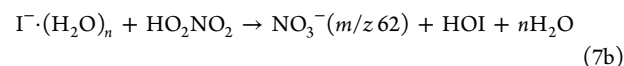
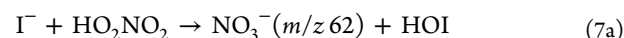
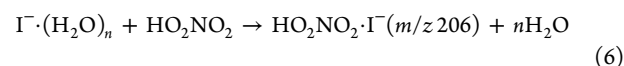
## EXPERIMENTAL METHODS

**Observations.** Observations were made in both winter (November 20, 2014 to January 2, 2015) and summer (June 10 to 7 July, 2015), on top of the Georgia Tech Ford ES&T building in metropolitan Atlanta (33°46' N, 84°24' W), 10 m above the ground, and 0.6 km to the west of the Interstate-75/85 highway. Concurrent measurements of HO<sub>2</sub>NO<sub>2</sub>, NO, NO<sub>2</sub>, CO, O<sub>3</sub>, and N<sub>2</sub>O<sub>5</sub> (winter only) and meteorological parameters were performed (Table 1). All instruments for gas phase species were housed in a laboratory on the top floor with inlets extending through the roof to sample ambient air. Basic meteorological parameters were measured by a weather station mounted on the roof of the building. Data were reported in local time, UTC−5 h in winter, and UTC−4 h in summer.

**CIMS Measurements of HO<sub>2</sub>NO<sub>2</sub>.** The general CIMS instrumentation and methodology have been discussed by Huey.<sup>41</sup> Due to instrument availability, we employed different CIMS instruments for each season. Both CIMS instruments had configurations similar to that described by Liao et al.<sup>29</sup> The main differences were inlet configurations and flows. The inlets were Teflon FEP tubing (O.D. = 1.22 cm and I. D. = 0.95 cm), 9 and 5 m long, respectively, in winter and summer observations. The inlet flows were 46 slpm (standard liter per minute, standard = 0 °C and 1013 hPa) and 41 slpm, respectively, of which roughly 1.5 slpm was sampled into CIMS and the rest was exhausted. These large flows resulted in relatively short residence times in the inlets, roughly 0.8 and 0.5 s, respectively. We also assessed the inlet loss of HO<sub>2</sub>NO<sub>2</sub> under the same conditions (i.e., flow rate and temperature) as the ambient measurement. HO<sub>2</sub>NO<sub>2</sub> was synthesized (Supporting Information), diluted, and flown through (1) the inlet used for sampling,

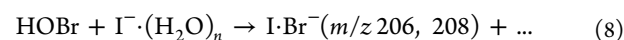
(2) a new, clean inlet of the same diameter and length as the sampling inlet, and (3) a very short (~30 cm), clean inlet of the same diameter. HO<sub>2</sub>NO<sub>2</sub> signals were found to be the same for all cases, and thus, we concluded that the wall loss of HO<sub>2</sub>NO<sub>2</sub> was minimal.

The detection of HO<sub>2</sub>NO<sub>2</sub> using I<sup>−</sup>(H<sub>2</sub>O)<sub>*n*</sub> CIMS was similar to that described by Veres et al.<sup>27</sup> Briefly, HO<sub>2</sub>NO<sub>2</sub> is ionized by I<sup>−</sup>(H<sub>2</sub>O)<sub>*n*</sub> and detected on two channels: HO<sub>2</sub>NO<sub>2</sub>·I<sup>−</sup> and NO<sub>3</sub><sup>−</sup>.



We retrieved HO<sub>2</sub>NO<sub>2</sub> concentrations based on the HO<sub>2</sub>NO<sub>2</sub>·I<sup>−</sup> channel. We found that the NO<sub>3</sub><sup>−</sup> channel is the major channel with a branching ratio >90%. However, it is known to be nonspecific for HO<sub>2</sub>NO<sub>2</sub>. It can be formed via reactions of I<sup>−</sup>(H<sub>2</sub>O)<sub>*n*</sub> and several other species (N<sub>2</sub>O<sub>5</sub>, NO<sub>3</sub>, BrONO<sub>2</sub>, ClONO<sub>2</sub>).<sup>37,41–43</sup> An assessment of the NO<sub>3</sub><sup>−</sup> channel is given in the Supporting Information.

As we used a quadrupole mass spectrometer of unit mass resolution, the detection at the HO<sub>2</sub>NO<sub>2</sub>·I<sup>−</sup> channel can be interfered with by I<sup>−79</sup>Br<sup>−</sup> of the same *m/z* in a bromine-rich environment.<sup>30</sup>



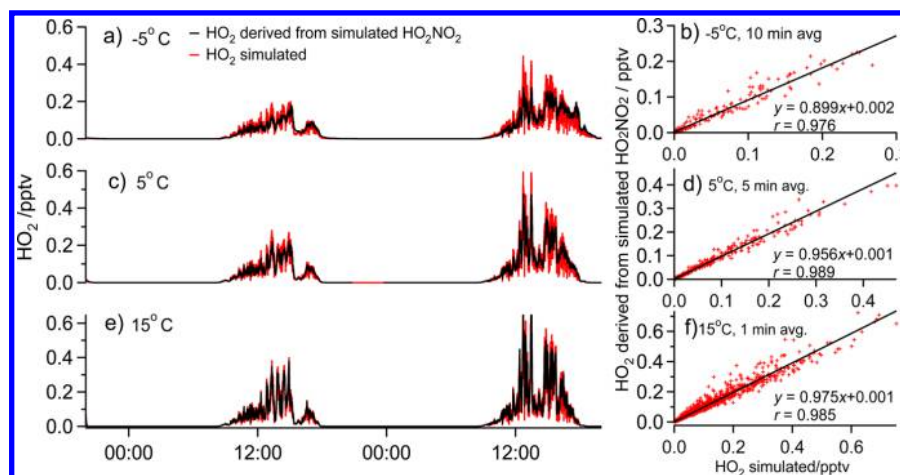
Since the sensitivities to isotopes are proportional to their abundances, this interference can be corrected by tracking the isotope I<sup>81</sup>Br<sup>−</sup> at *m/z* 208. Nitryl chloride (ClNO<sub>2</sub>) is detected as ClNO<sub>2</sub>·I<sup>−</sup> at *m/z* 208 (<sup>35</sup>ClNO<sub>2</sub>·I<sup>−</sup>) and 210 (<sup>37</sup>ClNO<sub>2</sub>·I<sup>−</sup>).<sup>33</sup> If ClNO<sub>2</sub> is present, the <sup>35</sup>ClNO<sub>2</sub>·I<sup>−</sup> signal at *m/z* 208 needs to be subtracted to obtain the I<sup>81</sup>Br<sup>−</sup> signal, and thus, the <sup>37</sup>ClNO<sub>2</sub>·I<sup>−</sup> signal at *m/z* 210 needs to be tracked. In winter, signals at *m/z* 208 and 210 were found to only originate from ClNO<sub>2</sub>, based on their correlation and the isotope ratios, which ruled out the I<sup>79</sup>Br<sup>−</sup> interference. In summer, no ClNO<sub>2</sub>·I<sup>−</sup> nor I·Br<sup>−</sup> was found at *m/z* 208 or 210.

We had previously observed that high NO<sub>2</sub> (500 ppbv) can interfere with HO<sub>2</sub>NO<sub>2</sub> detection using I<sup>−</sup> with a high resolution time-of-flight CIMS.<sup>44</sup> For this reason, NO<sub>2</sub> (up to 86 ppbv) was added to the inlet to test for possible interferences to the

Table 2. Summary of the Time-Dependent Model

	species/parameters
constrained	$T$ , <sup>a</sup> $p$ , <sup>b</sup> $O_3$ , <sup>c</sup> $NO$ , <sup>c</sup> $NO_2$ , <sup>c</sup> $H_2O(v)$ , <sup>d</sup> $CO$ , <sup>d</sup> $CH_4$ , <sup>d</sup> $C_2H_6$ , <sup>d</sup> $C_3H_8$ , <sup>d</sup> $CH_2O$ , <sup>d</sup> $CH_3C(O)CH_3$ , <sup>d</sup> $j$ values, <sup>e</sup>
time dependent	$HO_2NO_2$ , <sup>f</sup> $N_2O_5$ , <sup>f</sup> $HONO$ , <sup>f</sup> $PAN$ <sup>f</sup>
steady state	$OH$ , $HO_2$ , $CH_3O_2$ , $CH_3C(O)O_2$ , $NO_3$

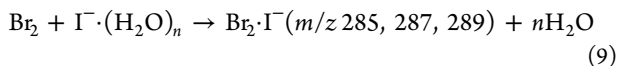
<sup>a</sup>Fixed to different values for each run. <sup>b</sup>Fixed to be 100 kPa. <sup>c</sup>Constrained by measurements. <sup>d</sup>Fixed concentrations:  $[H_2O(v)] = 8000$  ppmv,  $[CO] = 100$  ppbv,  $[CH_4] = 1.8$  ppmv,  $[C_2H_6] = 1$  ppbv,  $[C_3H_8] = 500$  pptv,  $[CH_2O] = 1$  ppbv, and  $[CH_3C(O)CH_3] = 2$  ppbv. <sup>e</sup>Estimated using the NCAR TUV model (<https://www2.acom.ucar.edu/modeling/tropospheric-ultraviolet-and-visible-tuv-radiation-model>). <sup>f</sup>Initial concentrations for time-dependent species:  $[HO_2NO_2]_0 = 2$  pptv,  $[N_2O_5]_0 = 50$  pptv,  $[HONO]_0 = 1$  ppbv, and  $[PAN]_0 = 1$  ppbv.



**Figure 1.**  $HO_2$ ,  $HO_2NO_2$ , and  $HO_2$  were simulated using the time-dependent model for a period of 24 h, with  $\Delta t = 10$  s, at  $-5$ ,  $5$ , and  $15$  °C. (a, c, e) Time series of derived  $HO_2$  from simulated  $HO_2NO_2$  is plotted in red, and simulated  $HO_2$  is plotted in black. (b, d, f) The derived and simulated  $HO_2$  levels were averaged on different time scales and compared.

ambient measurement. However, no significant interferences due to  $NO_2$  were found on either the  $HO_2NO_2 \cdot I^-$  or the  $NO_3^-$  channels.

A two-step indirect calibration was applied, as a portable  $HO_2NO_2$  calibration source has not been developed. A known amount of  $Br_2$  from a permeation tube (KIN-TEK) was periodically added into the inlet to track the sensitivity variation.<sup>28,29</sup>



$Br_2$  is a convenient proxy as the CIMS instrument has similar sensitivities to  $HO_2NO_2$  and  $Br_2$  on the cluster channels. The  $Br_2$  emission rate was determined every 2 or 3 days by a spectrophotometric method.<sup>45,46</sup> The relative sensitivity of  $HO_2NO_2$  to  $Br_2$  was determined in the laboratory by comparing signals of known amounts of  $HO_2NO_2$  and  $Br_2$ .  $HO_2NO_2$  was quantified by measuring the resulting  $NO_2$  of  $HO_2NO_2$  thermal decomposition with a cavity attenuated phase-shift spectroscopy<sup>47</sup> (CAPS, Aerodyne) (Supporting Information). The sensitivities to  $HO_2NO_2$  ( $HO_2NO_2 \cdot I^-$  channel) were  $\sim 25$  and  $\sim 15$  Hz·pptv<sup>-1</sup>, respectively, in both seasons. Commonly, we have estimated the limit of detection (LOD) based on the standard deviation of the variations of background signals.<sup>29,48</sup> In this work, we defined the LOD as twice the standard deviation ( $2\sigma$ ) of the background variation. The LODs of  $HO_2NO_2$  for both seasons were 0.2 and 0.6 pptv, respectively, based on a 1 min integration time. The uncertainties for  $HO_2NO_2$  measurement were 22% in winter and 16% in summer, mostly from calibration.

**Determination of  $HO_2$ .**  $HO_2NO_2$  observations were combined with measurements of  $NO_2$  and temperature to derive  $HO_2$  levels from eq 5.  $HO_2NO_2$  data were excluded during

precipitation events, as they perturbed the equilibrium, or if they were below the LOD. All  $NO_2$  data were well above the LOD. We applied different averaging times to derive  $HO_2$  at different temperatures. Temperature ranged from  $-4.7$  to  $24.5$  °C during the winter observation and  $18.8$  to  $37.6$  °C during the summer observation. The lifetime of  $HO_2NO_2$  and thus the time scale to approach equilibrium varied. Also,  $NO_x$  emitted by traffic (especially traffic on the highway nearby) fluctuated, which can perturb the  $HO_2NO_2$  equilibrium on short time scales.

We used a time-dependent model<sup>21</sup> to determine the appropriate averaging times. The model setup highlighted the influence of temperature and  $NO_x$  (Table 2). The chemical scheme is detailed in the Supporting Information. Model simulations were constrained to observed  $NO_x$  and  $O_3$  from December 1, 2014, 20:00, to December 3, 2014, 20:00, when variations of  $NO_x$  were especially pronounced. Model simulations were performed for this period for a set of assumed temperatures.  $HO_2$  was simulated directly from the model and also derived from model simulated  $HO_2NO_2$  using eq 5. Both model simulated and experimentally derived  $HO_2$  concentrations were averaged on different time scales and compared. The comparisons are shown in Figure 1. At  $-5$  °C, 10 min averages of derived and simulated  $HO_2$  agree within 10% (slope = 0.899,  $r = 0.976$ ); at  $5$  °C, 5 min averages of derived and simulated  $HO_2$  agree within 5% (slope = 0.955,  $r = 0.988$ ); at  $15$  °C, 1 min averages of derived and simulated  $HO_2$  agree within 3% (slope = 0.975,  $r = 0.985$ ). Thus, we select the average time of 10 min for  $-5$  °C  $\leq T < 5$  °C, 5 min for  $5$  °C  $\leq T < 15$  °C, and 1 min for  $T \geq 15$  °C.

The uncertainty for  $HO_2$  derived from  $HO_2NO_2$  observations has multiple sources. The equilibrium constant  $K$  in eq 5 has a 30% uncertainty in the temperature range of our



Table 3. Species and Parameters for the Steady-State Model Prediction of HO<sub>2</sub>

	species/parameters
constrained (measured)	T, p, O <sub>3</sub> , H <sub>2</sub> O(v), CO, NO, NO <sub>2</sub> , HO <sub>2</sub> NO <sub>2</sub> , <sup>a</sup> N <sub>2</sub> O, <sup>b</sup>
constrained (modeled/assumed)	j values, <sup>c</sup> CH <sub>4</sub> , <sup>d</sup> C <sub>2</sub> H <sub>6</sub> , <sup>e</sup> C <sub>3</sub> H <sub>8</sub> , <sup>e</sup> C <sub>2</sub> H <sub>4</sub> , <sup>e</sup> C <sub>3</sub> H <sub>6</sub> , <sup>e</sup> CH <sub>2</sub> O, <sup>e</sup> CH <sub>3</sub> OH, <sup>e</sup> CH <sub>3</sub> CHO, <sup>e</sup> PAN, <sup>e</sup> CH <sub>3</sub> C(O)CH <sub>3</sub> , <sup>e</sup> CH <sub>3</sub> C(O)CH <sub>2</sub> OH, <sup>e</sup> CH <sub>3</sub> C(O)CHO, <sup>e</sup> isoprene, <sup>e</sup> CH <sub>3</sub> C(O)CH=CH <sub>2</sub> (MVK), <sup>e</sup> CH <sub>2</sub> =C(CH <sub>3</sub> )CHO (MACR), <sup>e</sup> benzene, <sup>e</sup> toluene, <sup>e</sup> HONO, <sup>e</sup> H <sub>2</sub> O <sub>2</sub> <sup>e</sup>
predicted by the model	OH, HO <sub>2</sub> , CH <sub>3</sub> O <sub>2</sub> , C <sub>2</sub> H <sub>5</sub> O <sub>2</sub> , C <sub>2</sub> H <sub>4</sub> (OH)O <sub>2</sub> , CH <sub>3</sub> C(O)O <sub>2</sub> , CH <sub>3</sub> C(O)CH <sub>2</sub> O <sub>2</sub> , Isop(OH)O <sub>2</sub> , <sup>f</sup> MOHO <sub>2</sub> , <sup>g</sup> MCO <sub>3</sub> , <sup>h</sup> ArO <sub>2</sub> , <sup>i</sup> R(NO <sub>3</sub> )O <sub>2</sub> , <sup>j</sup> NO <sub>3</sub> , CH <sub>3</sub> OOH

<sup>a</sup>Predicted by the steady-state model in Run\_1; constrained by measurements in Run\_2 (see text). <sup>b</sup>Constrained by measurement in the winter observation; estimated using the REAM model in the summer observation. <sup>c</sup>Estimated using the NCAR TUV model, scaled using the measured UV radiation. <sup>d</sup>CH<sub>4</sub> was fixed to be 1.8 ppmv. <sup>e</sup>Estimated using the REAM model. <sup>f</sup>Hydroxylperoxy radicals derived from the OH-initiated oxidation of isoprene. <sup>g</sup>Hydroxylperoxy radicals derived from the OH-initiated oxidations of MACR/MVK. <sup>h</sup>CH<sub>2</sub>=C(CH<sub>3</sub>)C(O)O<sub>2</sub>. <sup>i</sup>Peroxy radicals derived from the oxidation of benzene and toluene. <sup>j</sup>Peroxy radicals derived from the NO<sub>3</sub>-alkene addition reactions.

observations<sup>13,15,49,50</sup> and is the major uncertainty source. The uncertainties of the HO<sub>2</sub>NO<sub>2</sub> and NO<sub>2</sub> measurements are listed in Table 1. Combining these effects gives a total uncertainty of 39% in the winter observation and 35% in the summer observation for derived HO<sub>2</sub>.

**Model Calculation of HO<sub>2</sub>.** A steady-state model was used to estimate HO<sub>2</sub> to compare to HO<sub>2</sub> derived from observations. The model was developed on the basis of previous work in our laboratory.<sup>20,21,51</sup> The model setup is listed in Table 3. The chemistry scheme is listed in the Supporting Information. The model included detected species and meteorological parameters. VOC, HONO, and H<sub>2</sub>O<sub>2</sub> were not measured. We estimated basic VOC (C<sub>2</sub>H<sub>6</sub>, C<sub>3</sub>H<sub>8</sub>, C<sub>2</sub>H<sub>4</sub>, C<sub>3</sub>H<sub>6</sub>, CH<sub>2</sub>O, CH<sub>3</sub>OH, CH<sub>3</sub>CHO, peroxyacetyl nitrate (PAN), CH<sub>3</sub>C(O)CH<sub>3</sub>, CH<sub>3</sub>C(O)CH<sub>2</sub>OH, CH<sub>3</sub>C(O)CHO, isoprene, methacrolein (MACR), methyl vinyl ketone (MVK), benzene, toluene), HONO, and H<sub>2</sub>O<sub>2</sub> using the 3-D Regional chEmical and trAnsport Model (REAM). Details of REAM can be found in previous studies.<sup>52–56</sup> Average levels of key species given by REAM in both seasons are listed in Table 4. j values were calculated using

Table 4. Average Concentrations of Key Species Predicted by REAM<sup>a</sup>

species	winter 2014	summer 2015
CH <sub>2</sub> O	1.3	5.0
CH <sub>3</sub> C(O)CH <sub>3</sub>	2.1	2.3
HONO	0.76	0.30
toluene	0.83	0.41
isoprene	0.08	2.0

<sup>a</sup>unit: ppbv.

the NCAR TUV model (<https://www2.acom.ucar.edu/modeling/tropospheric-ultraviolet-and-visible-tuv-radiation-model>) and scaled using measured solar UV radiation. The model was run twice. Run\_1 was not constrained by measured HO<sub>2</sub>NO<sub>2</sub> but calculated HO<sub>2</sub>NO<sub>2</sub> to be in steady state. Run\_2 included HO<sub>2</sub>NO<sub>2</sub> as a constraint. Run\_2 effectively makes HO<sub>2</sub>NO<sub>2</sub> a HO<sub>x</sub> source in the model and allows some compensation for sources that could be missing in the model.

## RESULTS AND DISCUSSION

Figures S7 and S8 show time series of HO<sub>2</sub>NO<sub>2</sub>, NO<sub>2</sub>, NO, temperature, solar radiation, and derived HO<sub>2</sub>. In winter, HO<sub>2</sub>NO<sub>2</sub> ranged from the LOD to 21.1 pptv (average 3.7 pptv), NO<sub>2</sub> ranged from 2.7 to 60.1 ppbv (average 17.0 ppbv), and derived HO<sub>2</sub> was up to 3.5 pptv (daytime average 0.17 pptv). In summer, HO<sub>2</sub>NO<sub>2</sub> ranged from the LOD to 90.3 pptv (average 11.7 pptv),

NO<sub>2</sub> ranged from 1.4 to 64.6 ppbv (average 11.3 ppbv), and derived HO<sub>2</sub> was up to 89.3 pptv (daytime average 10.4 pptv). Figure 2 shows mean diurnal profiles of HO<sub>2</sub>. Peak values of

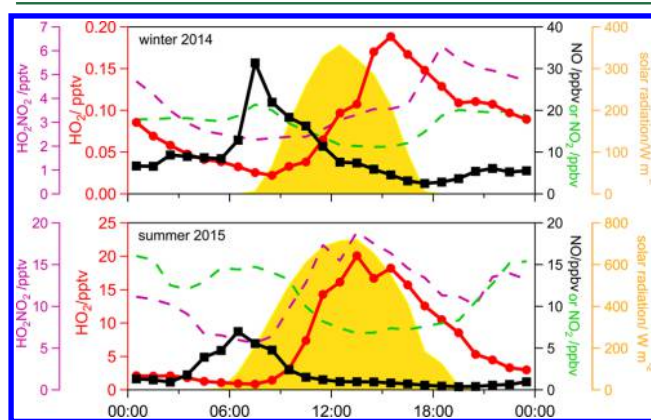


Figure 2. Diurnal profiles of derived HO<sub>2</sub>, NO, and solar radiation in both winter and summer observations. NO suppressed HO<sub>2</sub> during morning rush hours. Diurnal profiles of HO<sub>2</sub>NO<sub>2</sub> and NO<sub>2</sub> are also plotted in dotted lines.

HO<sub>2</sub> profiles were 0.19 pptv in winter and 20.0 pptv in summer. These HO<sub>2</sub> levels are close to those observed in the same seasons at polluted sites.<sup>6,57–59</sup> In the following sections, we look at two test cases and model comparisons to determine if the derived HO<sub>2</sub> is reasonable.

**Suppression of HO<sub>2</sub> at High NO.** HO<sub>2</sub> can be efficiently suppressed at high NO due to the fast reaction of HO<sub>2</sub> with NO. The most obvious suppression of HO<sub>2</sub> occurred during morning rush hour, as seen in Figure 2. In both seasons, the peak of NO in the morning depressed HO<sub>2</sub>, resulting in asymmetric diurnal profiles, although the radical production is expected to be symmetric as it follows the solar radiation. A similar behavior was observed in urban areas by Holland et al.<sup>60</sup> and Martinez et al.<sup>57</sup>

Low HO<sub>2</sub> associated with high NO was oftentimes observed in both seasons, during the day and night. Figure 3 displays two examples of the suppression of HO<sub>2</sub> at high NO peaks. On December 26 and 27, 2014, HO<sub>2</sub> decreased accordingly at high NO multiple times (20:00–23:00, 03:00–06:00, 07:40–08:30, 09:00–14:00). On June 20, 2015, HO<sub>2</sub> was around 1 pptv during 17:00–18:00 while NO was high up to 12 ppbv; then, HO<sub>2</sub> recovered to ~6 pptv after NO dropped to ~1 ppbv. In all, our derived HO<sub>2</sub> responded to NO accordingly. This indicates that there are no significant interferences when NO<sub>x</sub> is high that lead to an obvious over estimation

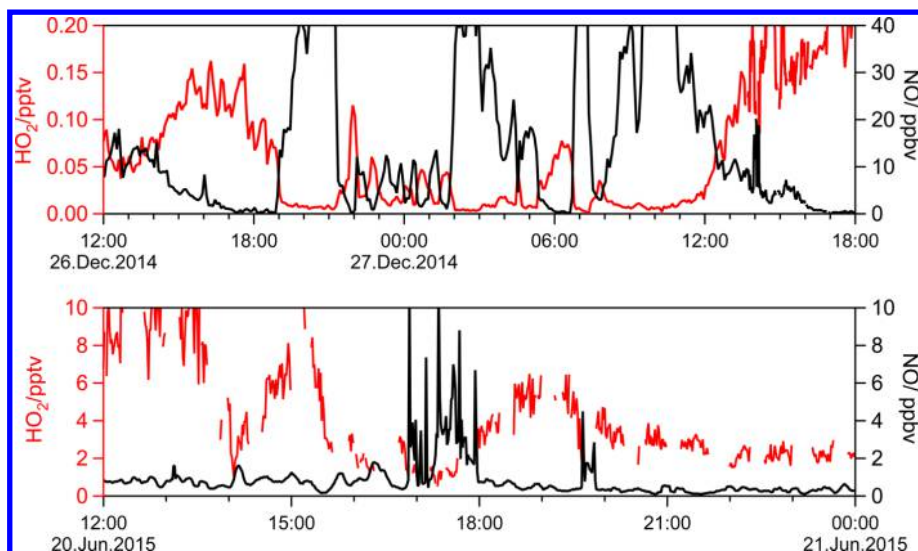


Figure 3. Examples of the suppression of HO<sub>2</sub> at high NO.

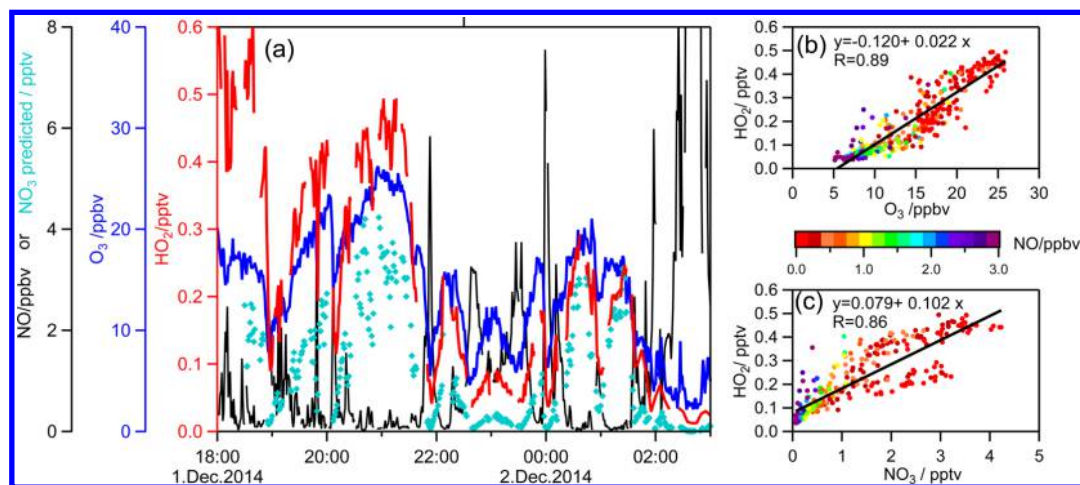


Figure 4. Correlation between HO<sub>2</sub> and O<sub>3</sub>/NO<sub>3</sub> during 18:30–01:30 on the night of December 1 to 2, 2014.

of HO<sub>2</sub>. However, very low levels of HO<sub>2</sub> were still observed at high NO.

**Relationship between HO<sub>2</sub> and O<sub>3</sub>/NO<sub>3</sub> at Night.** Nighttime HO<sub>x</sub> has been reported in several locations.<sup>3,61</sup> Nighttime HO<sub>2</sub> is largely produced by ozonolysis and oxidation of VOC by NO<sub>3</sub>.<sup>1,2,4,5,62–64</sup> Here, we examine the correlation between nighttime HO<sub>2</sub> and O<sub>3</sub> in both seasons and the correlation between nighttime HO<sub>2</sub> and NO<sub>3</sub> in winter when N<sub>2</sub>O<sub>5</sub> was measured, which allowed for an estimate of [NO<sub>3</sub>]:

$$[\text{NO}_3] \approx \frac{k_{\text{N}_2\text{O}_5}[\text{N}_2\text{O}_5] + k_{\text{NO}_2+\text{O}_3}[\text{NO}_2][\text{O}_3]}{k_{\text{NO}_3+\text{NO}_2}[\text{NO}_2] + k_{\text{NO}_3+\text{NO}}[\text{NO}]} \quad (10)$$

where  $k_{\text{N}_2\text{O}_5}$  is the unimolecular thermal decomposition rate of N<sub>2</sub>O<sub>5</sub>.

HO<sub>2</sub> was moderately correlated to O<sub>3</sub> at night for the both seasons.  $r(\text{HO}_2\text{--O}_3) = 0.68$  in the winter observation and 0.67 in summer. The correlation with NO<sub>3</sub> was found to be less,  $r(\text{HO}_2\text{--NO}_3) = 0.27$  in winter. However, if a filter of NO < 1 ppbv is applied, as high NO can suppress HO<sub>2</sub>, O<sub>3</sub>, and NO<sub>3</sub>, then  $r(\text{HO}_2\text{--O}_3) = 0.59$  in winter and 0.47 in summer and  $r(\text{HO}_2\text{--NO}_3) = 0.67$  in winter.

Strong correlations between HO<sub>2</sub> and O<sub>3</sub>/NO<sub>3</sub> were also observed on individual nights. For instance, as seen in Figure 4, during 18:30–01:30 on the night of December 1 to 2, 2014, most of the time NO < 1 ppbv, HO<sub>2</sub>, O<sub>3</sub> and NO<sub>3</sub> all showed similar temporal patterns,  $r(\text{HO}_2\text{--O}_3) = 0.89$ , and  $r(\text{HO}_2\text{--NO}_3) = 0.86$ . Also, as seen in Figure 5, during 20:00–05:00 on the night of June 30 to July 1, 2015, NO was mostly < 1 ppbv and  $r(\text{HO}_2\text{--O}_3) = 0.90$ . These strong correlations suggest that O<sub>3</sub> and NO<sub>3</sub> were driving the production of HO<sub>2</sub> on these nights.

**Comparison of Observation-Derived and Model-Predicted HO<sub>2</sub>.** Figure 6 compares the derived HO<sub>2</sub> with model predicted HO<sub>2</sub> (Run\_1 and Run\_2). (See Figures S7 and S8 for day-to-day variations of derived/model predicted HO<sub>2</sub> and other species.) As the model was not constrained by measured VOC, H<sub>2</sub>O<sub>2</sub>, HONO, and *j* values, discrepancies between model simulated and derived HO<sub>2</sub> are expected. Nevertheless, Run\_1 still captures the features of HO<sub>2</sub>. The average model-to-observed ratio (M/O) of Run\_1 is 1.27 ( $r = 0.54$ ) for winter and 0.70 ( $r = 0.80$ ) for summer. For the summer observations, Run\_1 tends to underestimate midday (e.g., June 15–18, 2015) and nighttime HO<sub>2</sub>, which suggests that more sources of HO<sub>2</sub> may exist than are included in the model.

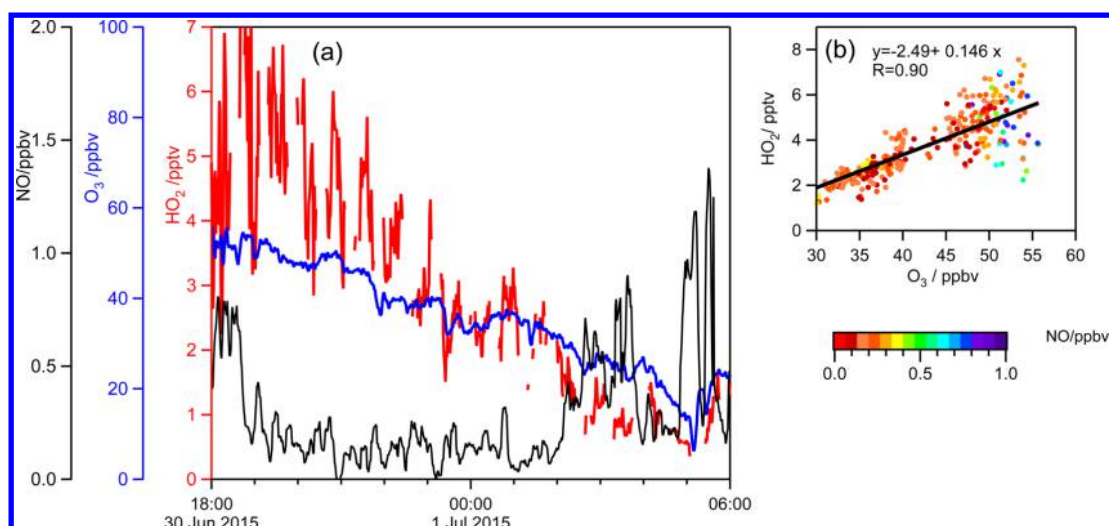


Figure 5. Correlation between  $\text{HO}_2$  and  $\text{O}_3$  during 20:00–05:00 on the night of June 30 to July 1, 2015.

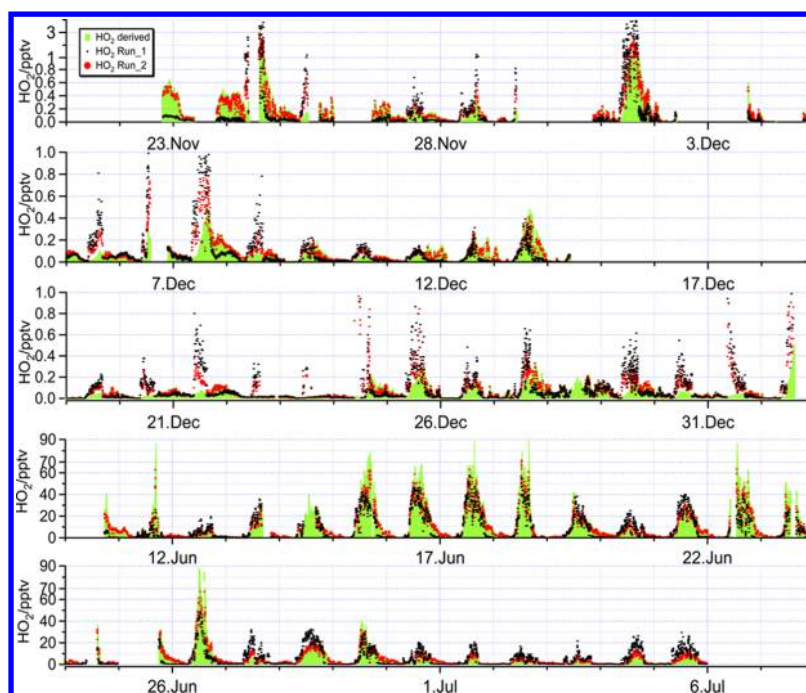


Figure 6. Comparison between the  $\text{HO}_2$  derived from observations to that predicted by the steady-state model. Run\_1 was not constrained by measured  $\text{HO}_2\text{NO}_2$ . Run\_2 was constrained by measured  $\text{HO}_2\text{NO}_2$ . Note that scales are different in different seasons.

With the constraint of  $\text{HO}_2\text{NO}_2$ , Run\_2 predicted  $\text{HO}_2$  with average  $M/O = 1.13$  ( $r = 0.77$ ) for winter and average  $M/O = 0.90$  ( $r = 0.97$ ) for summer. Run\_2 slightly improved  $M/O$  for the winter season and improved  $M/O$  for the summer season. Both model runs predicted peak midday  $\text{HO}_2$  levels on several winter days, such as December 21–24, that were significantly higher than the observations. One likely explanation for this is the overprediction of HONO levels by the REAM model. HONO has heterogeneous sources,<sup>65–67</sup> which are difficult to parametrize. The photolysis of HONO becomes a more important source of  $\text{HO}_x$  in winter, as other sources such as the photolysis of ozone contribute less. As a result, predicted HONO has a larger impact on  $\text{HO}_2$  prediction in winter. This also suggests the importance of better observational constraints in future studies. Nevertheless, considering the few observational constraints available, both model runs are in reasonable agreement with the derived  $\text{HO}_2$ .

**Future Studies.** Our study shows the feasibility of deriving  $\text{HO}_2$  from  $\text{HO}_2\text{NO}_2$  measurements in warm environments. Future work should focus on: (1) model calculation of  $\text{HO}_2$  with a better constrained observational data set and (2) inter-comparison with  $\text{HO}_2$  measurement techniques such as laser-induced fluorescence (LIF).<sup>68</sup> In addition, a direct CIMS  $\text{HO}_2$  measurement technique using  $\text{Br}^-$  as the reagent ion has recently been developed.<sup>40</sup> We plan on further comparisons with this method as well.

## ■ ASSOCIATED CONTENT

### 📄 Supporting Information

The Supporting Information is available free of charge on the ACS Publications website at DOI: 10.1021/acs.est.6b05169.

Synthesis of  $\text{HO}_2\text{NO}_2$ ; ion chemistry of  $\text{HO}_2\text{NO}_2$  with  $\Gamma^-(\text{H}_2\text{O})_n$ ; calibration of  $\text{HO}_2\text{NO}_2$ ; assessment of the



CIMS  $\text{NO}_3^-$  channel; chemistry scheme for the time-dependent and steady-state models; time series of observed  $\text{HO}_2\text{NO}_2$ , temperature,  $\text{NO}$ ,  $\text{NO}_2$ , solar radiation, derived and model-predicted  $\text{HO}_2$  (PDF)

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### Notes

The authors declare no competing financial interest.

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## REFERENCES

- (1) Atkinson, R.; Carter, W. P. L. Kinetics and Mechanisms of the Gas-Phase Reactions of Ozone with Organic-Compounds under Atmospheric Conditions. *Chem. Rev.* **1984**, *84* (5), 437–470.
- (2) Johnson, D.; Marston, G. The gas-phase ozonolysis of unsaturated volatile organic compounds in the troposphere. *Chem. Soc. Rev.* **2008**, *37*, 699–716.
- (3) Geyer, A.; Bachmann, K.; Hofzumahaus, A.; Holland, F.; Konrad, S.; Klupfel, T.; Patz, H. W.; Perner, D.; Mihelcic, D.; Schafer, H. J.; Volz-Thomas, A.; Platt, U. Nighttime formation of peroxy and hydroxyl radicals during the BERLIOZ campaign: Observations and modeling studies. *J. Geophys. Res.* **2003**, *108* (D4), 8249.
- (4) Paulson, S. E.; Orlando, J. J. The reactions of ozone with alkenes: An important source of  $\text{HO}_x$  in the boundary layer. *Geophys. Res. Lett.* **1996**, *23* (25), 3727–3730.
- (5) Wayne, R. P.; Barnes, I.; Biggs, P.; Burrows, J. P.; Canosamas, C. E.; Hjorth, J.; Lebras, G.; Moortgat, G. K.; Perner, D.; Poulet, G.; Restelli, G.; Sidebottom, H. The nitrate radical - physics, chemistry, and the atmosphere. *Atmos. Environ., Part A* **1991**, *25* (1), 1–203.
- (6) Kanaya, Y.; Cao, R. Q.; Akimoto, H.; Fukuda, M.; Komazaki, Y.; Yokouchi, Y.; Koike, M.; Tanimoto, H.; Takegawa, N.; Kondo, Y. Urban photochemistry in central Tokyo: 1. Observed and modeled OH and  $\text{HO}_2$  radical concentrations during the winter and summer of 2004. *J. Geophys. Res.* **2007**, *112* (D21), D21312.
- (7) Carlton, A. G.; Wiedinmyer, C.; Kroll, J. H. A review of Secondary Organic Aerosol (SOA) formation from isoprene. *Atmos. Chem. Phys.* **2009**, *9* (14), 4987–5005.
- (8) Jenkin, M. E. Modelling the formation and composition of secondary organic aerosol from  $\alpha$ - and  $\beta$ -pinene ozonolysis using MCM v3. *Atmos. Chem. Phys.* **2004**, *4*, 1741–1757.
- (9) Finlayson-Pitts, B. J.; Pitts, J. N., Jr. *Chemistry of the upper and lower atmosphere: theory, experiments, and applications*; Academic Press: San Diego, USA, 2000.
- (10) Jiménez, E.; Gierczak, T.; Stark, H.; Burkholder, J. B.; Ravishankara, A. R. Reaction of OH with  $\text{HO}_2\text{NO}_2$  (peroxynitric acid): Rate coefficients between 218 and 335 K and product yields at 298 K. *J. Phys. Chem. A* **2004**, *108* (7), 1139–1149.
- (11) Levine, S. Z.; Uselman, W. M.; Chan, W. H.; Calvert, J. G.; Shaw, J. H. Kinetics and mechanism of  $\text{HO}_2 + \text{NO}_2$  reactions - significance of peroxyoxynitric acid formation in photochemical smog. *Chem. Phys. Lett.* **1977**, *48* (3), 528–535.
- (12) Niki, H.; Maker, P. D.; Savage, C. M.; Breitenbach, L. P. Fourier-transform IR spectroscopic observation of pernitric acid formed via  $\text{HOO} + \text{NO}_2 \rightarrow \text{HOONO}_2$ . *Chem. Phys. Lett.* **1977**, *45* (3), 564–566.
- (13) Gierczak, T.; Jiménez, E.; Riffault, V.; Burkholder, J. B.; Ravishankara, A. R. Thermal decomposition of  $\text{HO}_2\text{NO}_2$  (peroxynitric acid, PNA): Rate coefficient and determination of the enthalpy of formation. *J. Phys. Chem. A* **2005**, *109* (4), 586–596.
- (14) Slusher, D. L.; Huey, L. G.; Tanner, D. J.; Chen, G.; Davis, D. D.; Buhr, M.; Nowak, J. B.; Eisele, F. L.; Kosciuch, E.; Mauldin, R. L.; Lefer, B. L.; Shetter, R. E.; Dibb, J. E. Measurements of pernitric acid at the South Pole during ISCAT 2000. *Geophys. Res. Lett.* **2002**, *29* (21), 2011.
- (15) Burkholder, J. B.; Sander, S. P.; Abbatt, J. P. D.; Barker, J. R.; Huie, R. E.; Kolb, C. E.; Kurylo, M. J.; Orkin, V. L.; Wilmouth, D. M.; Wine, P. H. *Chemical kinetics and photochemical data for use in atmospheric studies, evaluation No. 18*; Jet Propulsion Laboratory: Pasadena, USA, 2015.
- (16) Talbot, R. W.; Dibb, J. E.; Scheuer, E. M.; Kondo, Y.; Koike, M.; Singh, H. B.; Salas, L. B.; Fukui, Y.; Ballenthin, J. O.; Meads, R. F.; Miller, T. M.; Hunton, D. E.; Viggiano, A. A.; Blake, D. R.; Blake, N. J.; Atlas, E.; Flocke, F.; Jacob, D. J.; Jaegle, L. Reactive nitrogen budget during the NASA SONEX mission. *Geophys. Res. Lett.* **1999**, *26* (20), 3057–3060.
- (17) Jaeglé, L.; Jacob, D. J.; Brune, W. H.; Wennberg, P. O. Chemistry of  $\text{HO}_x$  radicals in the upper troposphere. *Atmos. Environ.* **2001**, *35* (3), 469–489.
- (18) Chen, G.; Davis, D.; Crawford, J.; Hutterli, L. M.; Huey, L. G.; Slusher, D.; Mauldin, L.; Eisele, F.; Tanner, D.; Dibb, J.; Buhr, M.; McConnell, J.; Lefer, B.; Shetter, R.; Blake, D.; Song, C. H.; Lombardi, K.; Arnoldy, J. A reassessment of  $\text{HO}_x$  South Pole chemistry based on observations recorded during ISCAT 2000. *Atmos. Environ.* **2004**, *38* (32), 5451–5461.
- (19) Chen, G.; Davis, D.; Crawford, J.; Nowak, J. B.; Eisele, F.; Mauldin, R. L.; Tanner, D.; Buhr, M.; Shetter, R.; Lefer, B.; Arimoto, R.; Hogan, A.; Blake, D. An investigation of South Pole  $\text{HO}_x$  chemistry: Comparison of model results with ISCAT observations. *Geophys. Res. Lett.* **2001**, *28* (19), 3633–3636.
- (20) Sjostedt, S. Investigation of photochemistry at high latitudes: comparison of model predictions to measurements of short lived species. Ph. D. Thesis, Georgia Institute of Technology, Atlanta, USA, 2006.
- (21) Kim, S.; Huey, L. G.; Stickel, R. E.; Tanner, D. J.; Crawford, J. H.; Olson, J. R.; Chen, G.; Brune, W. H.; Ren, X.; Leshner, R.; Wooldridge, P. J.; Bertram, T. H.; Perring, A.; Cohen, R. C.; Lefer, B. L.; Shetter, R. E.; Avery, M.; Diskin, G.; Sokolik, I. Measurement of  $\text{HO}_2\text{NO}_2$  in the free troposphere during the intercontinental chemical transport experiment - North America 2004. *J. Geophys. Res.* **2007**, *112* (D12), D12S01.
- (22) Stiller, G. P.; von Clarmann, T.; Bruehl, C.; Fischer, H.; Funke, B.; Glatthor, N.; Grabowski, U.; Hopfner, M.; Jockel, P.; Kellmann, S.; Kiefer, M.; Linden, A.; Lopez-Puertas, M.; Tsidu, G. M.; Milz, M.; Steck, T.; Steil, B. Global distributions of  $\text{HO}_2\text{NO}_2$  as observed by the Michelson Interferometer for Passive Atmospheric Sounding (MIPAS). *J. Geophys. Res.* **2007**, *112* (D9), D09314.
- (23) Jones, A. E.; Brough, N.; Anderson, P. S.; Wolff, E. W.  $\text{HO}_2\text{NO}_2$  and  $\text{HNO}_3$  in the coastal Antarctic winter night: a "lab-in-the-field" experiment. *Atmos. Chem. Phys.* **2014**, *14* (21), 11843–11851.
- (24) Spencer, K. M.; McCabe, D. C.; Crouse, J. D.; Olson, J. R.; Crawford, J. H.; Weinheimer, A. J.; Knapp, D. J.; Montzka, D. D.; Cantrell, C. A.; Hornbrook, R. S.; Mauldin, R. L.; Wennberg, P. O. Inferring ozone production in an urban atmosphere using measure-

ments of peroxyxynitric acid. *Atmos. Chem. Phys.* **2009**, *9* (11), 3697–3707.

(25) Zhang, R. Y.; Leu, M. T.; Keyser, L. F. Heterogeneous chemistry of HO<sub>2</sub>NO<sub>2</sub> in liquid sulfuric acid. *J. Phys. Chem. A* **1997**, *101* (18), 3324–3330.

(26) Abida, O.; Mielke, L. H.; Osthoff, H. D. Observation of gas-phase peroxyxynitrous and peroxyxynitric acid during the photolysis of nitrate in acidified frozen solutions. *Chem. Phys. Lett.* **2011**, *511* (4–6), 187–192.

(27) Veres, P. R.; Roberts, J. M.; Wild, R. J.; Edwards, P. M.; Brown, S. S.; Bates, T. S.; Quinn, P. K.; Johnson, J. E.; Zamora, R. J.; de Gouw, J. Peroxyxynitric acid (HO<sub>2</sub>NO<sub>2</sub>) measurements during the UBWOS 2013 and 2014 studies using iodide ion chemical ionization mass spectrometry. *Atmos. Chem. Phys.* **2015**, *15* (14), 8101–8114.

(28) Neuman, J. A.; Nowak, J. B.; Huey, L. G.; Burkholder, J. B.; Dibb, J. E.; Holloway, J. S.; Liao, J.; Peischl, J.; Roberts, J. M.; Ryerson, T. B.; Scheuer, E.; Stark, H.; Sticker, R. E.; Tanner, D. J.; Weinheimer, A. Bromine measurements in ozone depleted air over the Arctic Ocean. *Atmos. Chem. Phys.* **2010**, *10* (14), 6503–6514.

(29) Liao, J.; Sihler, H.; Huey, L. G.; Neuman, J. A.; Tanner, D. J.; Friess, U.; Platt, U.; Flocke, F. M.; Orlando, J. J.; Shepson, P. B.; Beine, H. J.; Weinheimer, A. J.; Sjostedt, S. J.; Nowak, J. B.; Knapp, D. J.; Staebler, R. M.; Zheng, W.; Sander, R.; Hall, S. R.; Ullmann, K. A comparison of Arctic BrO measurements by chemical ionization mass spectrometry and long path-differential optical absorption spectroscopy. *J. Geophys. Res.* **2011**, *116* (D14), D00R02.

(30) Liao, J.; Huey, L. G.; Tanner, D. J.; Flocke, F. M.; Orlando, J. J.; Neuman, J. A.; Nowak, J. B.; Weinheimer, A. J.; Hall, S. R.; Smith, J. N.; Friedl, A.; Staebler, R. M.; Wang, Y.; Koo, J. H.; Cantrell, C. A.; Weibring, P.; Walega, J.; Knapp, D. J.; Shepson, P. B.; Stephens, C. R. Observations of inorganic bromine (HOBr, BrO, and Br<sub>2</sub>) speciation at Barrow, Alaska, in spring 2009. *J. Geophys. Res.-Atmos.* **2012**, *117* (D14), D00R16.

(31) Liao, J.; Huey, L. G.; Liu, Z.; Tanner, D. J.; Cantrell, C. A.; Orlando, J. J.; Flocke, F. M.; Shepson, P. B.; Weinheimer, A. J.; Hall, S. R.; Ullmann, K.; Beine, H. J.; Wang, Y. H.; Ingall, E. D.; Stephens, C. R.; Hornbrook, R. S.; Apel, E. C.; Riemer, D.; Friedl, A.; Mauldin, R. L.; Smith, J. N.; Staebler, R. M.; Neuman, J. A.; Nowak, J. B. High levels of molecular chlorine in the Arctic atmosphere. *Nat. Geosci.* **2014**, *7* (2), 91–94.

(32) Pratt, K. A.; Custard, K. D.; Shepson, P. B.; Douglas, T. A.; Pöhler, D.; General, S.; Zielcke, J.; Simpson, W. R.; Platt, U.; Tanner, D. J.; Huey, L. G.; Carlsen, M.; Stirm, B. H. Photochemical production of molecular bromine in Arctic surface snowpacks. *Nat. Geosci.* **2013**, *6* (5), 351–356.

(33) Kercher, J. P.; Riedel, T. P.; Thornton, J. A. Chlorine activation by N<sub>2</sub>O<sub>5</sub>: simultaneous, in situ detection of ClNO<sub>2</sub> and N<sub>2</sub>O<sub>5</sub> by chemical ionization mass spectrometry. *Atmos. Meas. Tech.* **2009**, *2* (1), 193–204.

(34) McNeill, V. F.; Patterson, J.; Wolfe, G. M.; Thornton, J. A. The effect of varying levels of surfactant on the reactive uptake of N<sub>2</sub>O<sub>5</sub> to aqueous aerosol. *Atmos. Chem. Phys.* **2006**, *6*, 1635–1644.

(35) Le Breton, M.; McGillen, M. R.; Muller, J. B. A.; Bacak, A.; Shallcross, D. E.; Xiao, P.; Huey, L. G.; Tanner, D.; Coe, H.; Percival, C. J. Airborne observations of formic acid using a chemical ionization mass spectrometer. *Atmos. Meas. Tech.* **2012**, *5* (12), 3029–3039.

(36) Le Breton, M.; Bacak, A.; Muller, J. B. A.; Xiao, P.; Shallcross, B. M. A.; Batt, R.; Cooke, M. C.; Shallcross, D. E.; Bauguitte, S. J. B.; Percival, C. J. Simultaneous airborne nitric acid and formic acid measurements using a chemical ionization mass spectrometer around the UK: Analysis of primary and secondary production pathways. *Atmos. Environ.* **2014**, *83*, 166–175.

(37) Thornton, J. A.; Braban, C. F.; Abbatt, J. P. D. N<sub>2</sub>O<sub>5</sub> hydrolysis on sub-micron organic aerosols: the effect of relative humidity, particle phase, and particle size. *Phys. Chem. Chem. Phys.* **2003**, *5* (20), 4593–4603.

(38) Woodward-Massey, R.; Taha, Y. M.; Moussa, S. G.; Osthoff, H. D. Comparison of negative-ion proton-transfer with iodide ion

chemical ionization mass spectrometry for quantification of isocyanic acid in ambient air. *Atmos. Environ.* **2014**, *98*, 693–703.

(39) Phillips, G. J.; Pouvesle, N.; Thieser, J.; Schuster, G.; Axinte, R.; Fischer, H.; Williams, J.; Lelieveld, J.; Crowley, J. N. Peroxyacetyl nitrate (PAN) and peroxyacetic acid (PAA) measurements by iodide chemical ionisation mass spectrometry: first analysis of results in the boreal forest and implications for the measurement of PAN fluxes. *Atmos. Chem. Phys.* **2013**, *13* (3), 1129–1139.

(40) Sanchez, J.; Tanner, D. J.; Chen, D.; Huey, L. G.; Ng, N. L. A new technique for the direct detection of HO<sub>2</sub> radicals using bromide chemical ionization mass spectrometry (Br-CIMS): initial characterization. *Atmos. Meas. Tech.* **2016**, *9*, 3851–3861.

(41) Huey, L. G. Measurement of trace atmospheric species by chemical ionization mass spectrometry: Speciation of reactive nitrogen and future directions. *Mass Spectrom. Rev.* **2007**, *26* (2), 166–184.

(42) Hanson, D. R.; Ravishankara, A. R. The Reaction Probabilities of ClONO<sub>2</sub> and N<sub>2</sub>O<sub>5</sub> on Polar Stratospheric Cloud Materials. *J. Geophys. Res.* **1991**, *96* (D3), 5081–5090.

(43) Slusher, D. L.; Huey, L. G.; Tanner, D. J.; Flocke, F. M.; Roberts, J. M. A thermal dissociation-chemical ionization mass spectrometry (TD-CIMS) technique for the simultaneous measurement of peroxyacyl nitrates and dinitrogen pentoxide. *J. Geophys. Res.* **2004**, *109* (D19), D19315.

(44) Sanchez, J. School of Chemical and Biochemical Engineering, Georgia Institute of Technology, Atlanta, GA, USA, personal communication, 2016.

(45) Wu, C.; Birky, M. M.; Hepler, L. G. Thermochemistry of some bromine and iodine species in aqueous solution. *J. Phys. Chem.* **1963**, *67* (6), 1202–1205.

(46) Kazantseva, N. N.; Ernepesova, A.; Khodjamamedov, A.; Geldyev, O. A.; Krumgalz, B. S. Spectrophotometric analysis of iodide oxidation by chlorine in highly mineralized solutions. *Anal. Chim. Acta* **2002**, *456* (1), 105–119.

(47) Kebabian, P. L.; Wood, E. C.; Herndon, S. C.; Freedman, A. A practical alternative to chemiluminescence-based detection of nitrogen dioxide: Cavity attenuated phase shift spectroscopy. *Environ. Sci. Technol.* **2008**, *42* (16), 6040–6045.

(48) Neuman, J. A.; Huey, L. G.; Dissly, R. W.; Fehsenfeld, F. C.; Flocke, F.; Holecck, J. C.; Holloway, J. S.; Hubler, G.; Jakoubek, R.; Nicks, D. K.; Parrish, D. D.; Ryerson, T. B.; Sueper, D. T.; Weinheimer, A. J. Fast-response airborne in situ measurements of HNO<sub>3</sub> during the Texas 2000 Air Quality Study. *J. Geophys. Res.* **2002**, *107* (D20), 4436.

(49) Christensen, L. E.; Okumura, M.; Sander, S. P.; Friedl, R. R.; Miller, C. E.; Sloan, J. J. Measurements of the rate constant of HO<sub>2</sub>+NO<sub>2</sub>+N<sub>2</sub>→HO<sub>2</sub>NO<sub>2</sub>+N<sub>2</sub> using near-infrared wavelength-modulation spectroscopy and UV-visible absorption spectroscopy. *J. Phys. Chem. A* **2004**, *108* (1), 80–91.

(50) Bacak, A.; Cooke, M. C.; Bardwell, M. W.; McGillen, M. R.; Archibald, A. T.; Huey, L. G.; Tanner, D.; Utembe, S. R.; Jenkin, M. E.; Derwent, R. G.; Shallcross, D. E.; Percival, C. J. Kinetics of the HO<sub>2</sub>+NO<sub>2</sub> Reaction: On the impact of new gas-phase kinetic data for the formation of HO<sub>2</sub>NO<sub>2</sub> on HO<sub>x</sub>, NO<sub>x</sub> and HO<sub>2</sub>NO<sub>2</sub> levels in the troposphere. *Atmos. Environ.* **2011**, *45* (35), 6414–6422.

(51) Liao, J.; Huey, L. G.; Tanner, D. J.; Brough, N.; Brooks, S.; Dibb, J. E.; Stutz, J.; Thomas, J. L.; Lefer, B.; Haman, C.; Gorham, K. Observations of hydroxyl and peroxy radicals and the impact of BrO at Summit, Greenland in 2007 and 2008. *Atmos. Chem. Phys.* **2011**, *11* (16), 8577–8591.

(52) Choi, Y.; Wang, Y. H.; Zeng, T.; Martin, R. V.; Kurosu, T. P.; Chance, K. Evidence of lightning NO<sub>x</sub> and convective transport of pollutants in satellite observations over North America. *Geophys. Res. Lett.* **2005**, *32* (2), L02805.

(53) Jing, P.; Cunnold, D.; Choi, Y.; Wang, Y. Summertime tropospheric ozone columns from Aura OMI/MLS measurements versus regional model results over the United States. *Geophys. Res. Lett.* **2006**, *33* (17), L17817.



- (54) Wang, Y. H.; Choi, Y. S.; Zeng, T.; Ridley, B.; Blake, N.; Blake, D.; Flocke, F. Late-spring increase of trans-Pacific pollution transport in the upper troposphere. *Geophys. Res. Lett.* **2006**, *33* (1), L01811.
- (55) Liu, Z.; Wang, Y. H.; Gu, D. S.; Zhao, C.; Huey, L. G.; Stickel, R.; Liao, J.; Shao, M.; Zhu, T.; Zeng, L. M.; Liu, S. C.; Chang, C. C.; Amoroso, A.; Costabile, F. Evidence of reactive aromatics as a major source of peroxy acetyl nitrate over China. *Environ. Sci. Technol.* **2010**, *44* (18), 7017–7022.
- (56) Zhao, C.; Wang, Y. H.; Yang, Q.; Fu, R.; Cunnold, D.; Choi, Y. Impact of East Asian summer monsoon on the air quality over China: View from space. *J. Geophys. Res.* **2010**, *115*, D09301.
- (57) Martinez, M.; Harder, H.; Kovacs, T. A.; Simpas, J. B.; Bassis, J.; Leshner, R.; Brune, W. H.; Frost, G. J.; Williams, E. J.; Stroud, C. A.; Jobson, B. T.; Roberts, J. M.; Hall, S. R.; Shetter, R. E.; Wert, B.; Fried, A.; Alicke, B.; Stutz, J.; Young, V. L.; White, A. B.; Zamora, R. J. OH and HO<sub>2</sub> concentrations, sources, and loss rates during the Southern Oxidants Study in Nashville, Tennessee, summer 1999. *J. Geophys. Res.-Atmos* **2003**, *108* (D19), 4617.
- (58) Ren, X. R.; Brune, W. H.; Mao, J. Q.; Mitchell, M. J.; Leshner, R. L.; Simpas, J. B.; Metcalf, A. R.; Schwab, J. J.; Cai, C. X.; Li, Y. Q.; Demerjian, K. L.; Felton, H. D.; Boynton, G.; Adams, A.; Perry, J.; He, Y.; Zhou, X. L.; Hou, J. Behavior of OH and HO<sub>2</sub> in the winter atmosphere in New York city. *Atmos. Environ.* **2006**, *40*, S252–S263.
- (59) Shirley, T. R.; Brune, W. H.; Ren, X.; Mao, J.; Leshner, R.; Cardenas, B.; Volkamer, R.; Molina, L. T.; Molina, M. J.; Lamb, B.; Velasco, E.; Jobson, T.; Alexander, M. Atmospheric oxidation in the Mexico City Metropolitan Area (MCMA) during April 2003. *Atmos. Chem. Phys.* **2006**, *6*, 2753–2765.
- (60) Holland, F.; Hofzumahaus, A.; Schafer, R.; Kraus, A.; Patz, H. W. Measurements of OH and HO<sub>2</sub> radical concentrations and photolysis frequencies during BERLIOZ. *J. Geophys. Res.* **2003**, *108* (D4), 8246.
- (61) Commane, R.; Floquet, C. F. A.; Ingham, T.; Stone, D.; Evans, M. J.; Heard, D. E. Observations of OH and HO<sub>2</sub> radicals over West Africa. *Atmos. Chem. Phys.* **2010**, *10* (18), 8783–8801.
- (62) Brown, S. S.; de Gouw, J. A.; Warneke, C.; Ryerson, T. B.; Dube, W. P.; Atlas, E.; Weber, R. J.; Peltier, R. E.; Neuman, J. A.; Roberts, J. M.; Swanson, A.; Flocke, F.; McKeen, S. A.; Brioude, J.; Sommariva, R.; Trainer, M.; Fehsenfeld, F. C.; Ravishankara, A. R. Nocturnal isoprene oxidation over the Northeast United States in summer and its impact on reactive nitrogen partitioning and secondary organic aerosol. *Atmos. Chem. Phys.* **2009**, *9* (9), 3027–3042.
- (63) Sommariva, R.; Osthoff, H. D.; Brown, S. S.; Bates, T. S.; Baynard, T.; Coffman, D.; de Gouw, J. A.; Goldan, P. D.; Kuster, W. C.; Lerner, B. M.; Stark, H.; Warneke, C.; Williams, E. J.; Fehsenfeld, F. C.; Ravishankara, A. R.; Trainer, M. Radicals in the marine boundary layer during NEAQS 2004: a model study of day-time and night-time sources and sinks. *Atmos. Chem. Phys.* **2009**, *9* (9), 3075–3093.
- (64) Sommariva, R.; Brown, S. S.; Roberts, J. M.; Brookes, D. M.; Parker, A. E.; Monks, P. S.; Bates, T. S.; Bon, D.; de Gouw, J. A.; Frost, G. J.; Gilman, J. B.; Goldan, P. D.; Herndon, S. C.; Kuster, W. C.; Lerner, B. M.; Osthoff, H. D.; Tucker, S. C.; Warneke, C.; Williams, E. J.; Zahniser, M. S. Ozone production in remote oceanic and industrial areas derived from ship based measurements of peroxy radicals during TexAQ5 2006. *Atmos. Chem. Phys.* **2011**, *11* (6), 2471–2485.
- (65) Zhou, X. L.; Gao, H. L.; He, Y.; Huang, G.; Bertman, S. B.; Civerolo, K.; Schwab, J. Nitric acid photolysis on surfaces in low-NO<sub>x</sub> environments: Significant atmospheric implications. *Geophys. Res. Lett.* **2003**, *30* (23), 2217.
- (66) Wang, S. H.; Ackermann, R.; Spicer, C. W.; Fast, J. D.; Schmeling, M.; Stutz, J. Atmospheric observations of enhanced NO<sub>2</sub>-HONO conversion on mineral dust particles. *Geophys. Res. Lett.* **2003**, *30* (11), 1595.
- (67) Liu, Z.; Wang, Y. H.; Costabile, F.; Amoroso, A.; Zhao, C.; Huey, L. G.; Stickel, R.; Liao, J.; Zhu, T. Evidence of aerosols as a media for rapid daytime HONO production over China. *Environ. Sci. Technol.* **2014**, *48* (24), 14386–14391.
- (68) Fuchs, H.; Holland, F.; Hofzumahaus, A. Measurement of tropospheric RO<sub>2</sub> and HO<sub>2</sub> radicals by a laser-induced fluorescence instrument. *Rev. Sci. Instrum.* **2008**, *79* (8), 084104.
- (69) Ryerson, T. B.; Williams, E. J.; Fehsenfeld, F. C. An efficient photolysis system for fast-response NO<sub>2</sub> measurements. *J. Geophys. Res.-Atmos* **2000**, *105* (D21), 26447–26461.